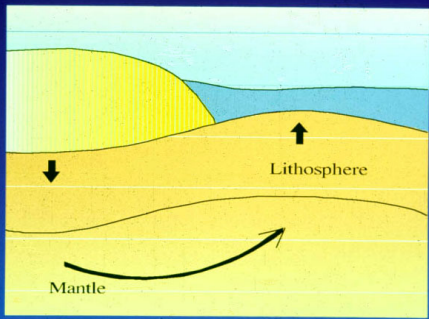


GIA signals in Geodetic Reference Frames

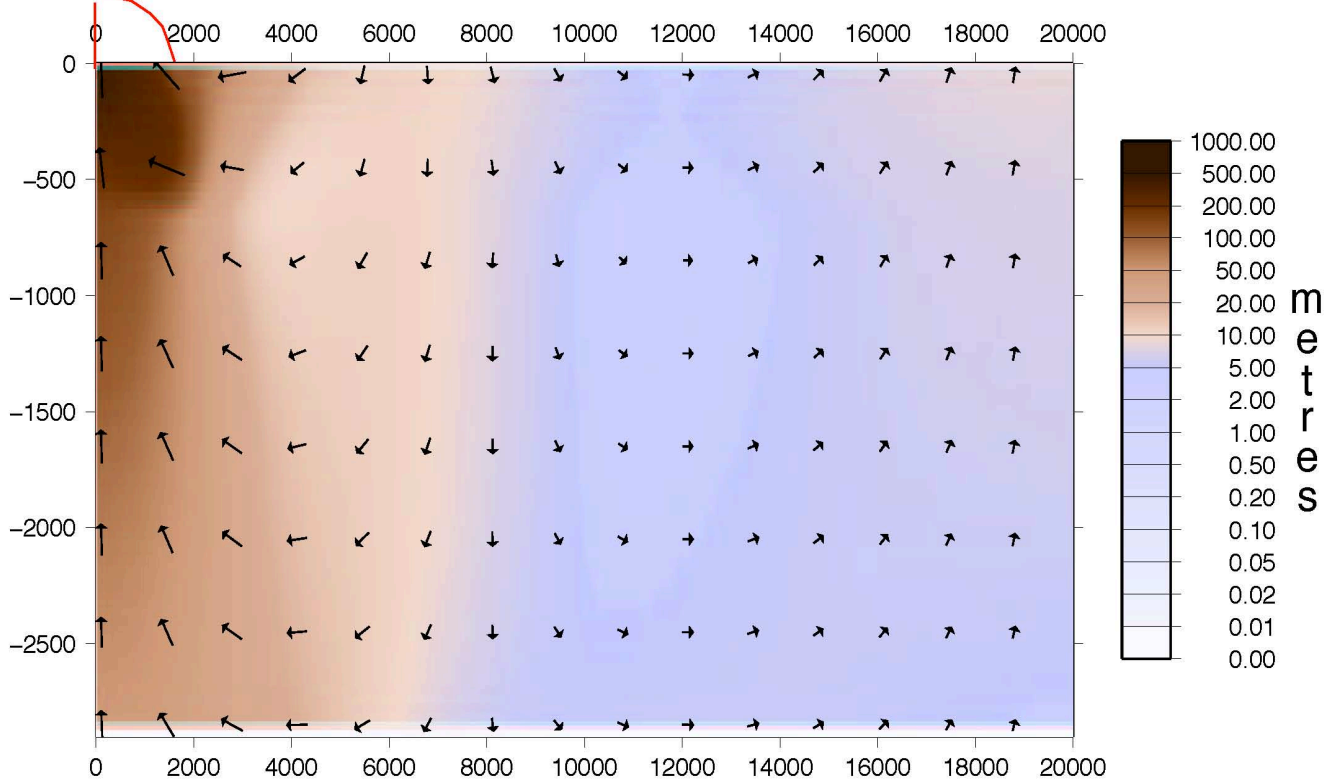
Kurt Lambeck, Anthony Purcell and H el ene Rouby



symmetric
parabolic
ice
load

Deformation due to an axisymmetric parabolic Heaviside load

Radius: 14.8 degrees Thickness: 3000 Time: 5 ka after unloading Earth model: ma4A



- (i) The deformation of the solid Earth in response to a single ice load change, is global.
- (ii) There is the associated water load change and changes in inertia tensor
- (iii) It is ongoing for $> 10^3$ years after the ice load is stabilized
- (iv) The signals will be indistinguishable from recent or present-day deglaciation from areas that also contributed to the past deglaciation.
- (v) To have effective predictive models we need to know the Earth rheology. Maxwell rheology? Internal phase boundaries? Lateral variability in viscosity and lithospheric thickness?
- (vi) Need the ice load history back to the LGM and beyond,
- (vii) Need high resolution models for the ocean bathymetry, particularly for former and present ice shelves.

Physical-Mathematical
model

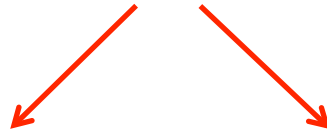
Starting earth-ice model



V_o, E_o



Observations



Far-field

near-field



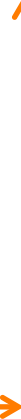
Function of V_o, E_o

Function of $\delta V_o, E_o$



Solve for improved
 V_o, E_o

Solve for improved
 $\delta V_o, E_o$



$$V_o = \sum \delta V_o ?$$

No ?

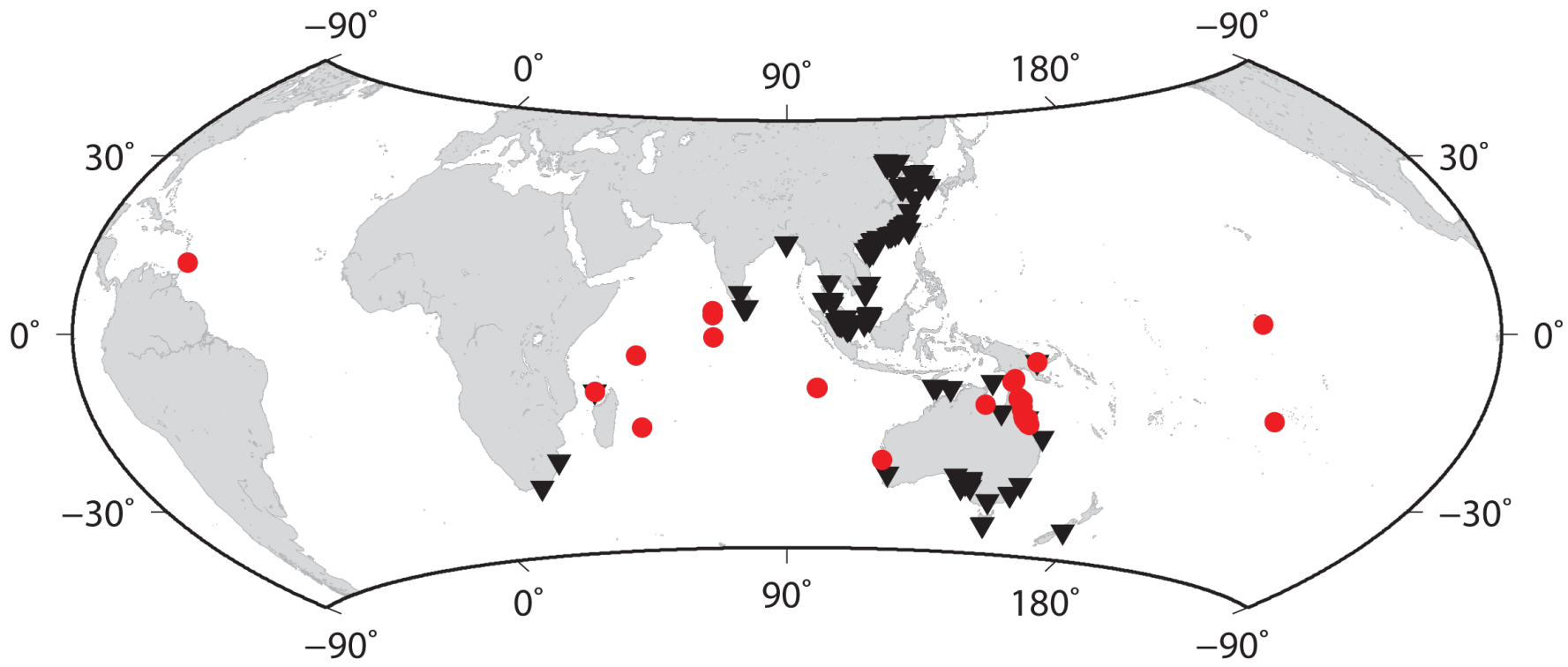
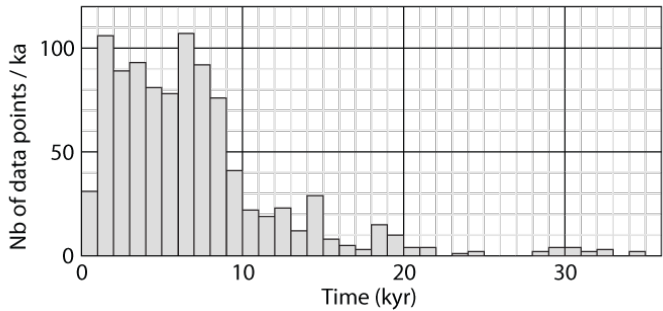
Yes ?

Then we have predictive model

Critical
Observations ?

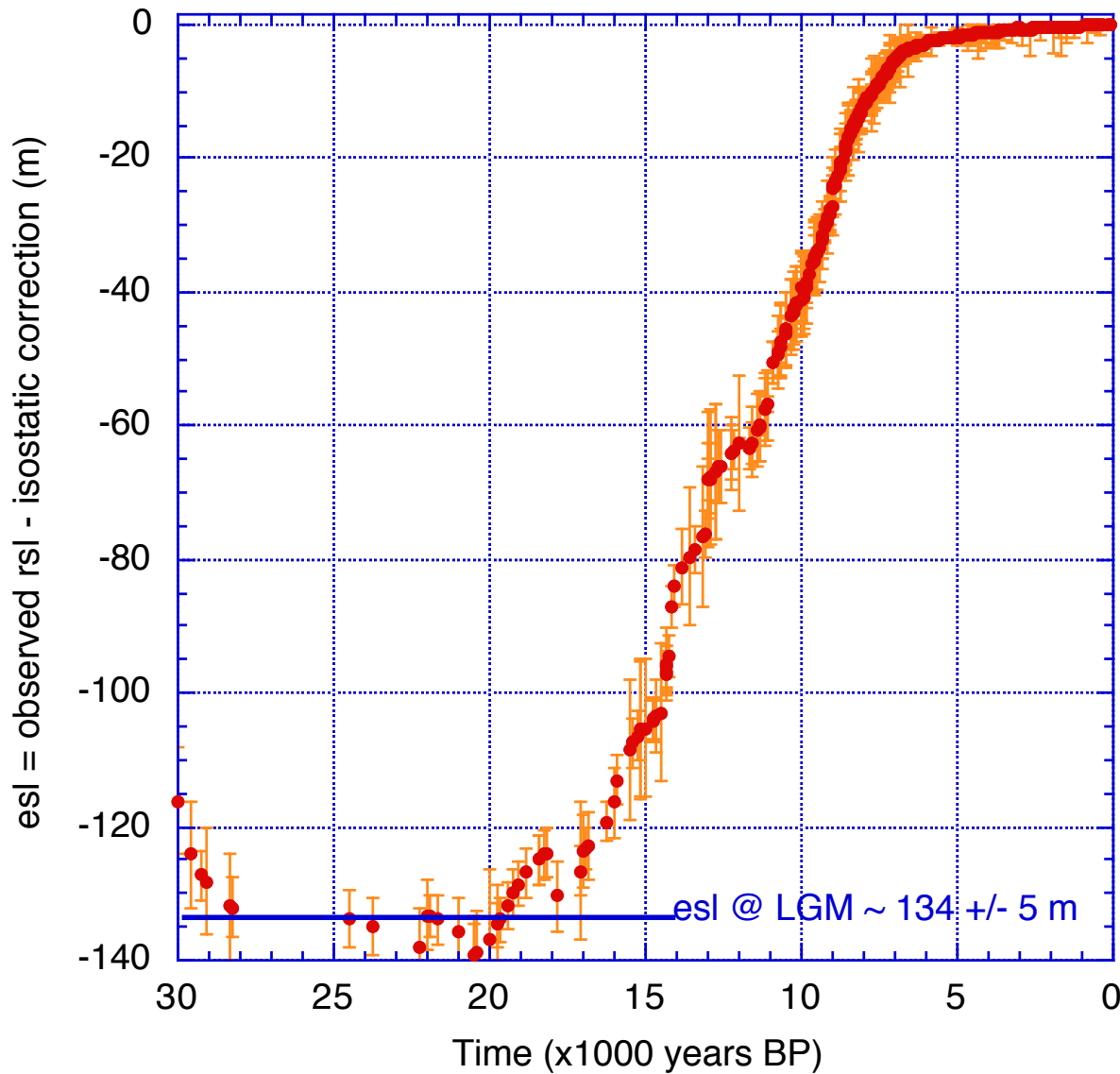
Critical
Observations ?

OBSERVATIONAL EVIDENCE



- Corals
- ▼ Sediments

Esl = ice volume equivalent sea level



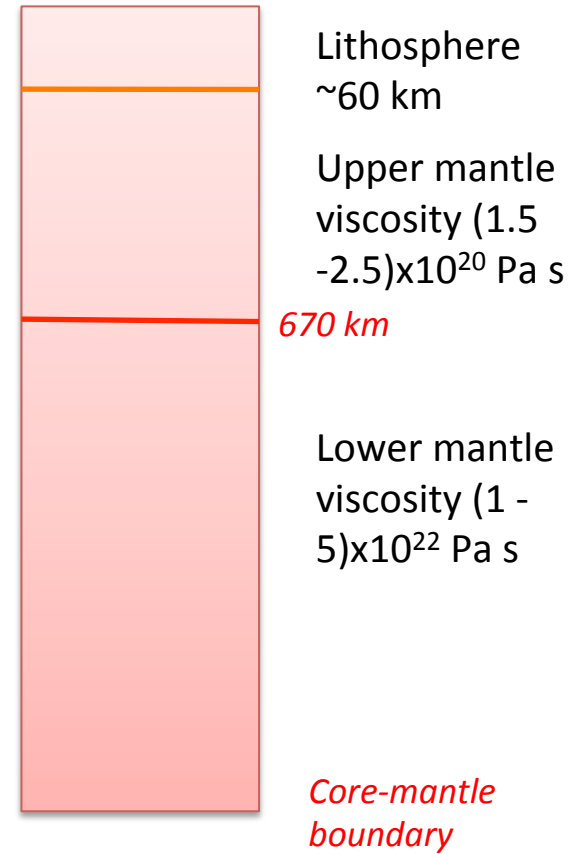
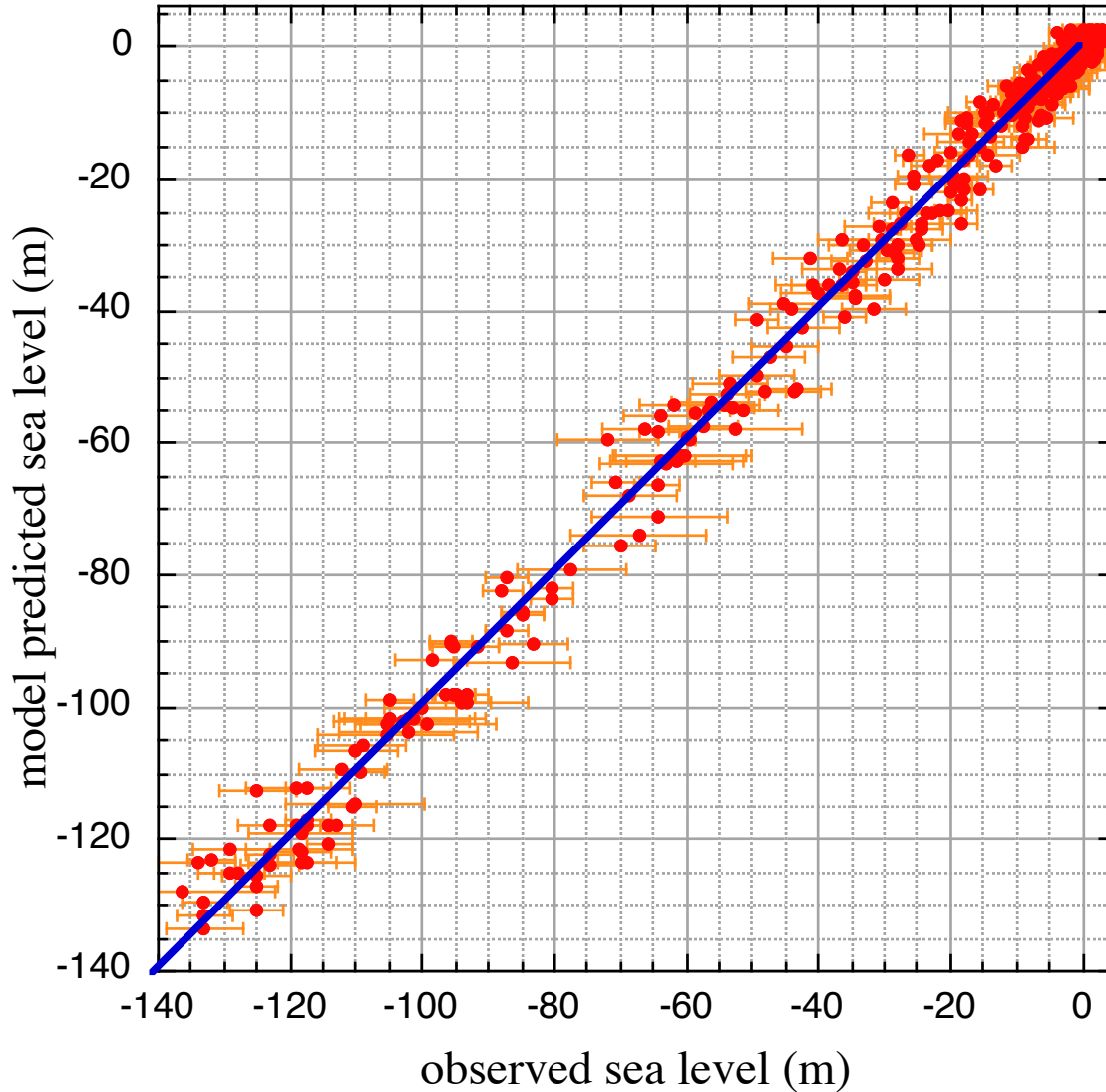
$$\Delta \zeta_{\text{esl}}(t) = -\frac{\rho_i}{\rho_o} \int_t^{\infty} \frac{1}{A_o(t)} \frac{dV_i}{dt} dt$$

$$\Delta \zeta_{\text{esl}} = \Delta \zeta_{\text{obs}} - (\Delta \zeta_I + \Delta \zeta_T)$$

134 m esl = 52×10^6
km³ grounded ice

SOLUTION (FF9-9)MtaB WITH FFC

19_16_16_12.RtaB.far#56102D

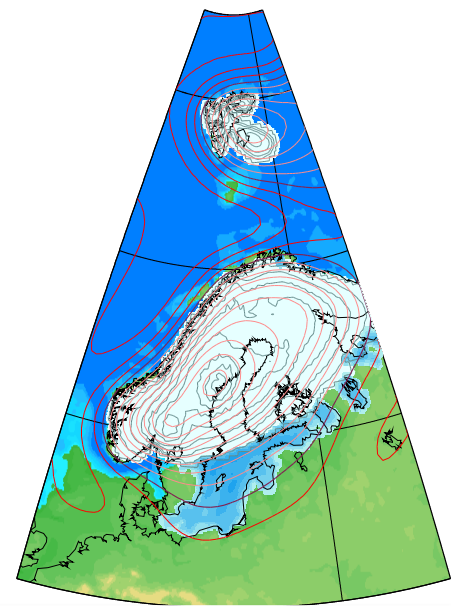
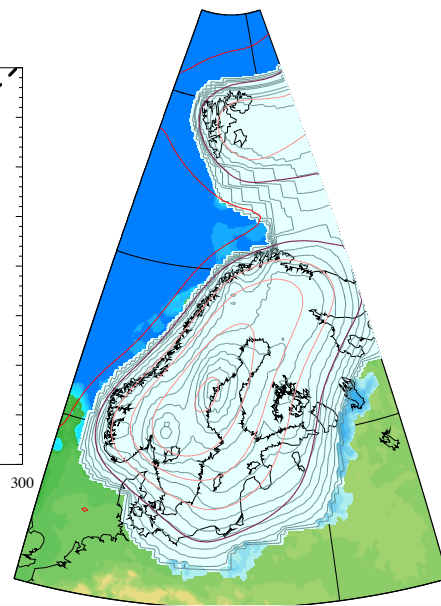
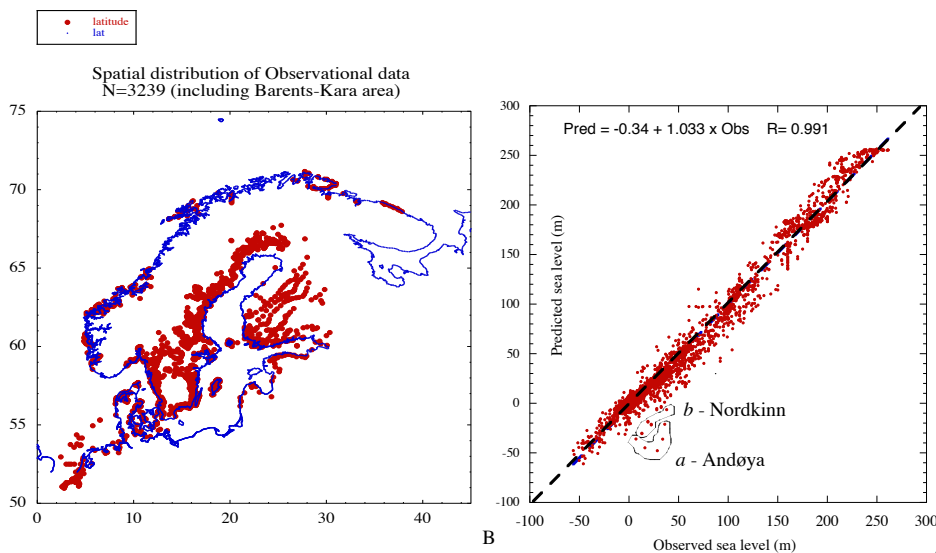


Earth response parameters for Fennoscandia

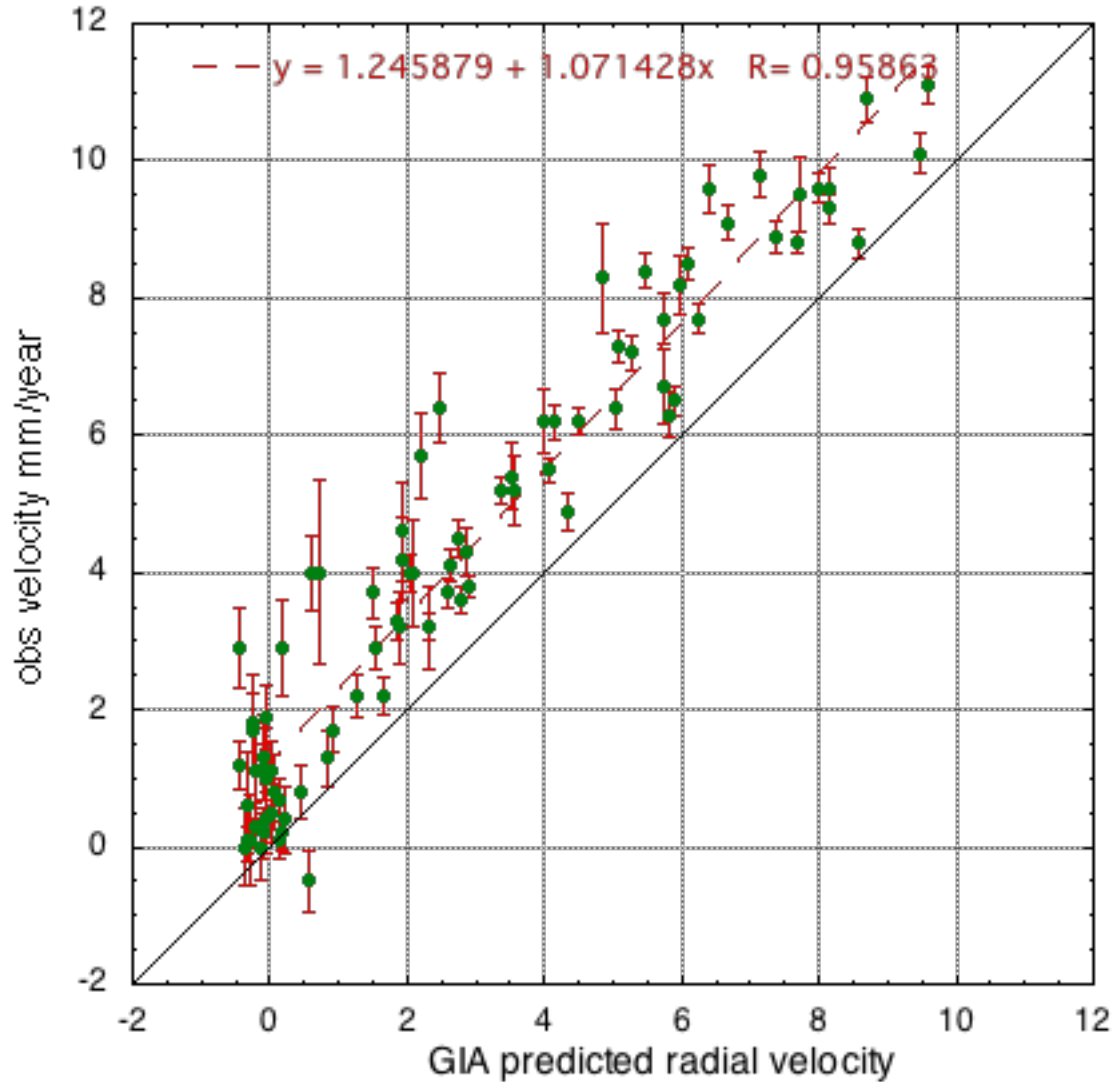
Litho- sphere 80 km	Upper mantle (3-4) $\times 10^{20}$ Pa s	Lower mantle (5-20) $\times 10^{22}$ Pa s
---------------------------	---	--

	H (km)	η_{um} ($\times 10^{20}$ Pa s)	η_{lm} ($\times 10^{21}$ Pa s)
Geological data ^{1,2}	80-90	3.0-4.0	7-30
Tide gauge data ^{1,3}	80-100	4.0-4.5	≥ 10
GPS data ⁴	93-110	3.4-5.0	7-13

¹Lambeck and Purcell (2003); ²Lambeck et al., (2008); ³Lambeck and Ekman (1998); ⁴Zhao et al. (2012).



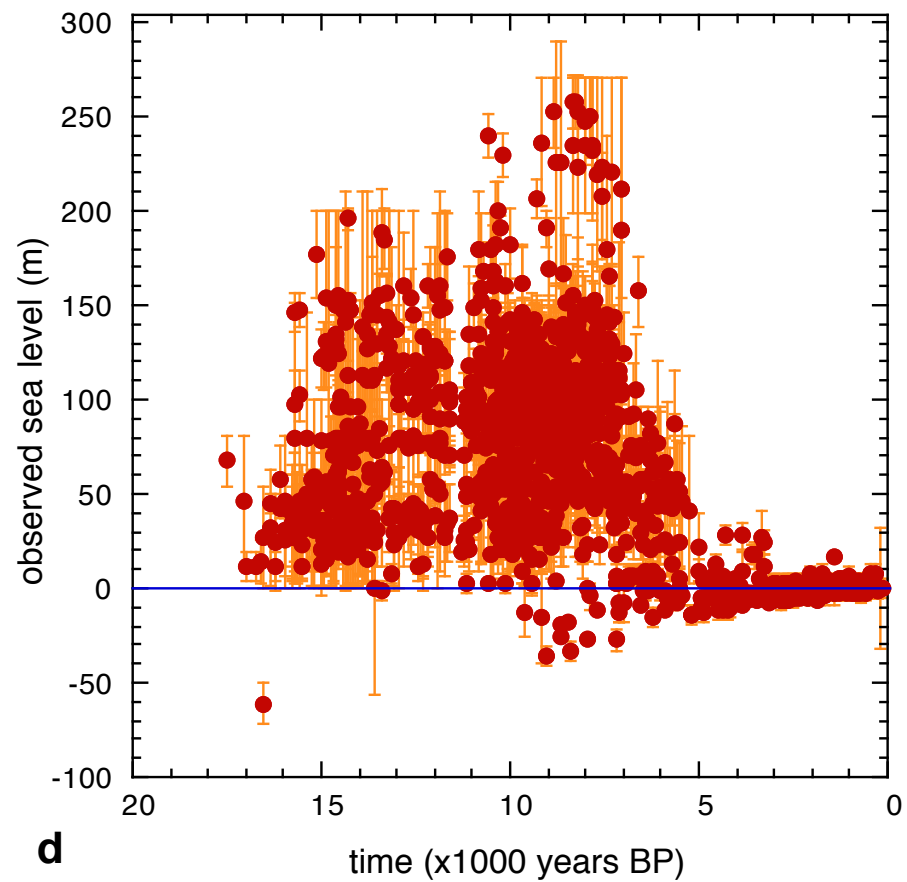
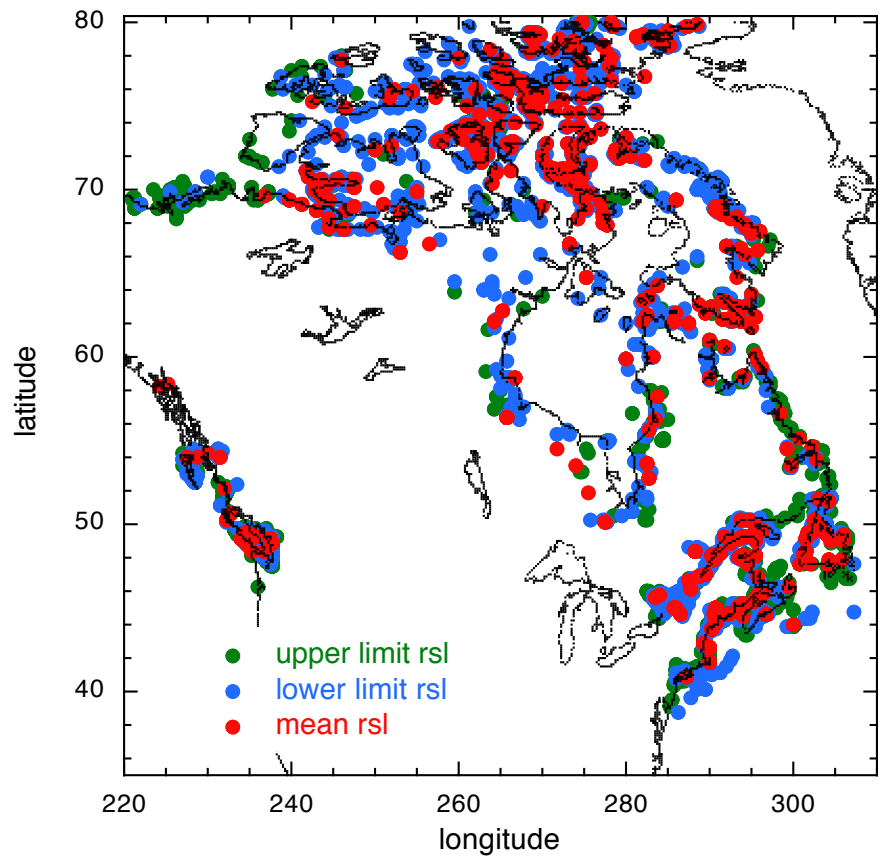
vert.K33_j1b_WS9_6.ma4B
Bifrost-ma4B new Calsea

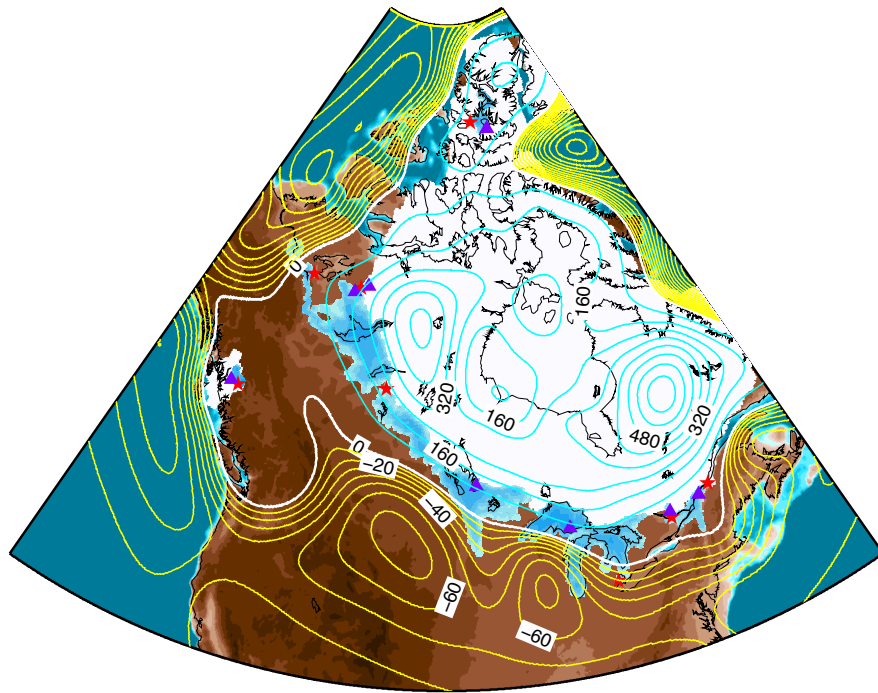


Observed radial velocity:
Lidberg 2009. (ITRF 2005)

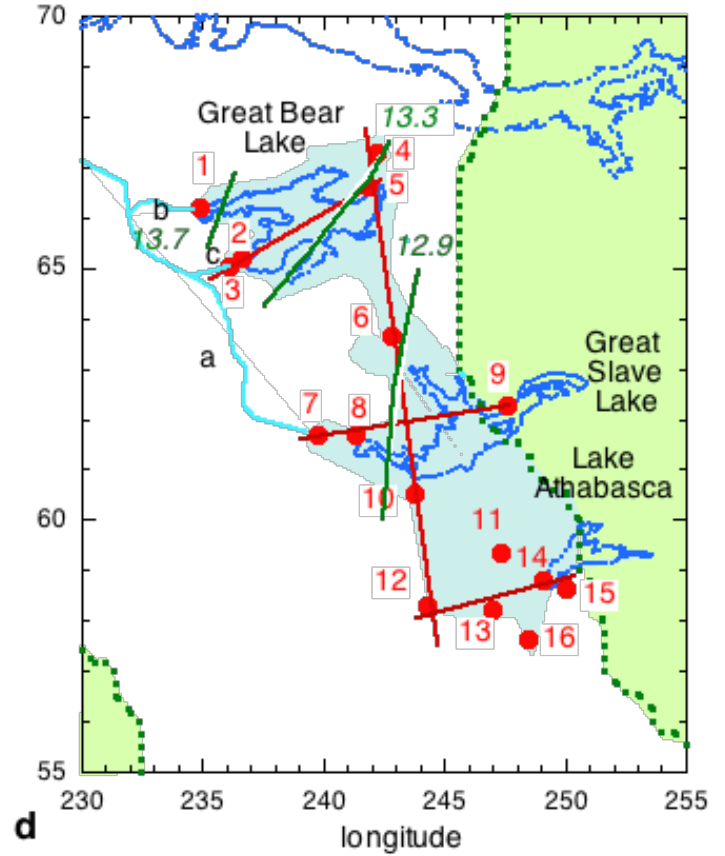
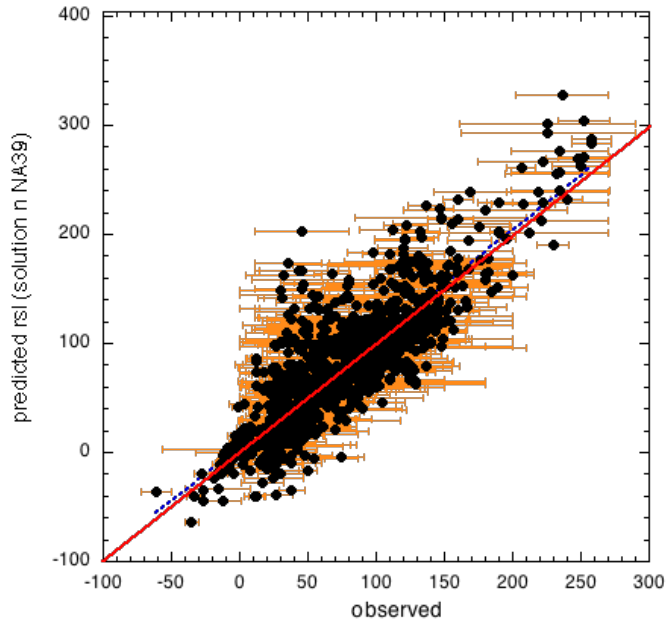
Predicted radial velocity:
GIA model Lambeck et al. 2010

Noted a systematic offset with
observed rates > predicted rates.

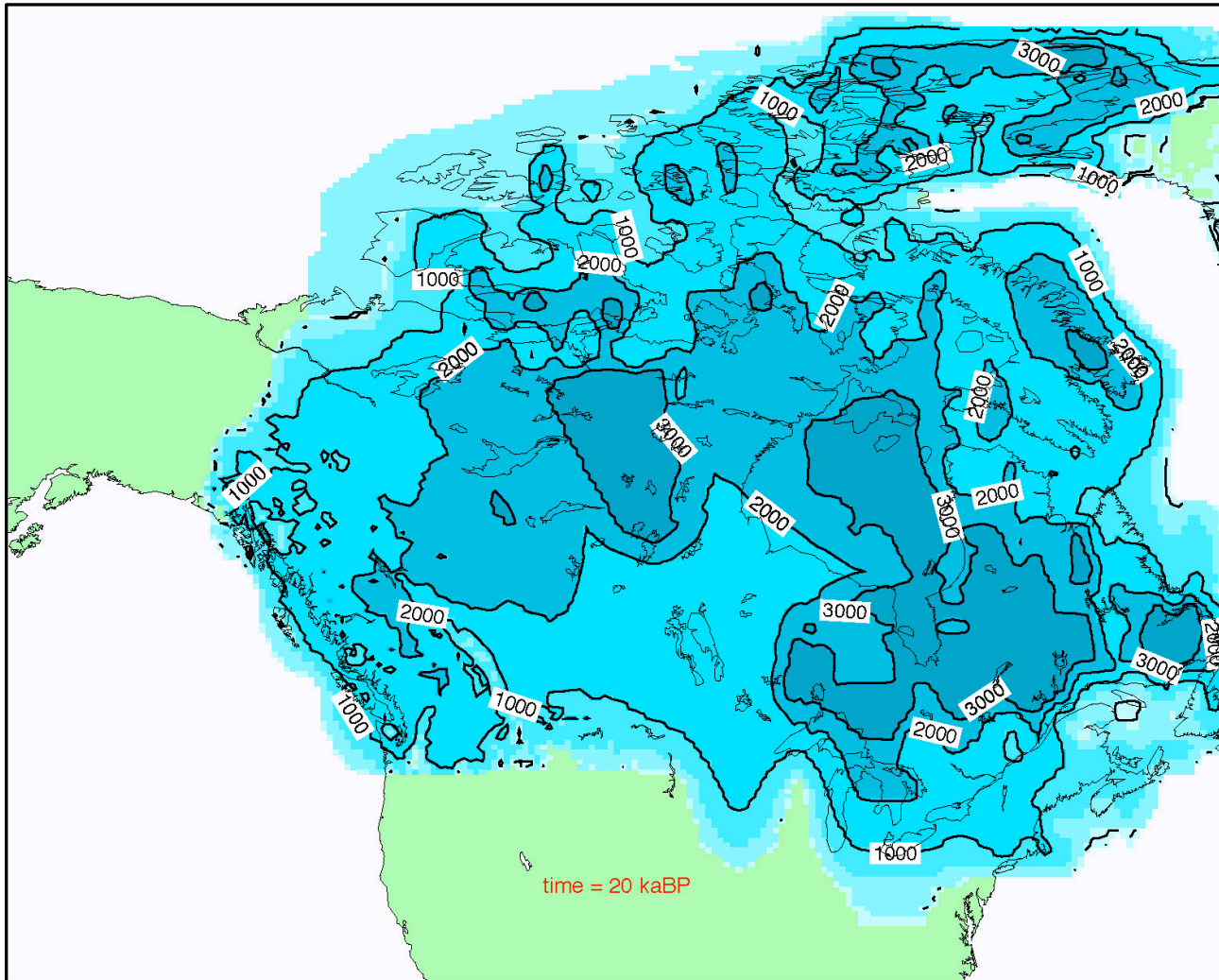
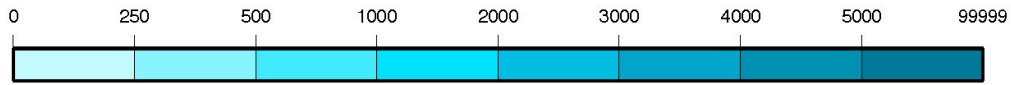




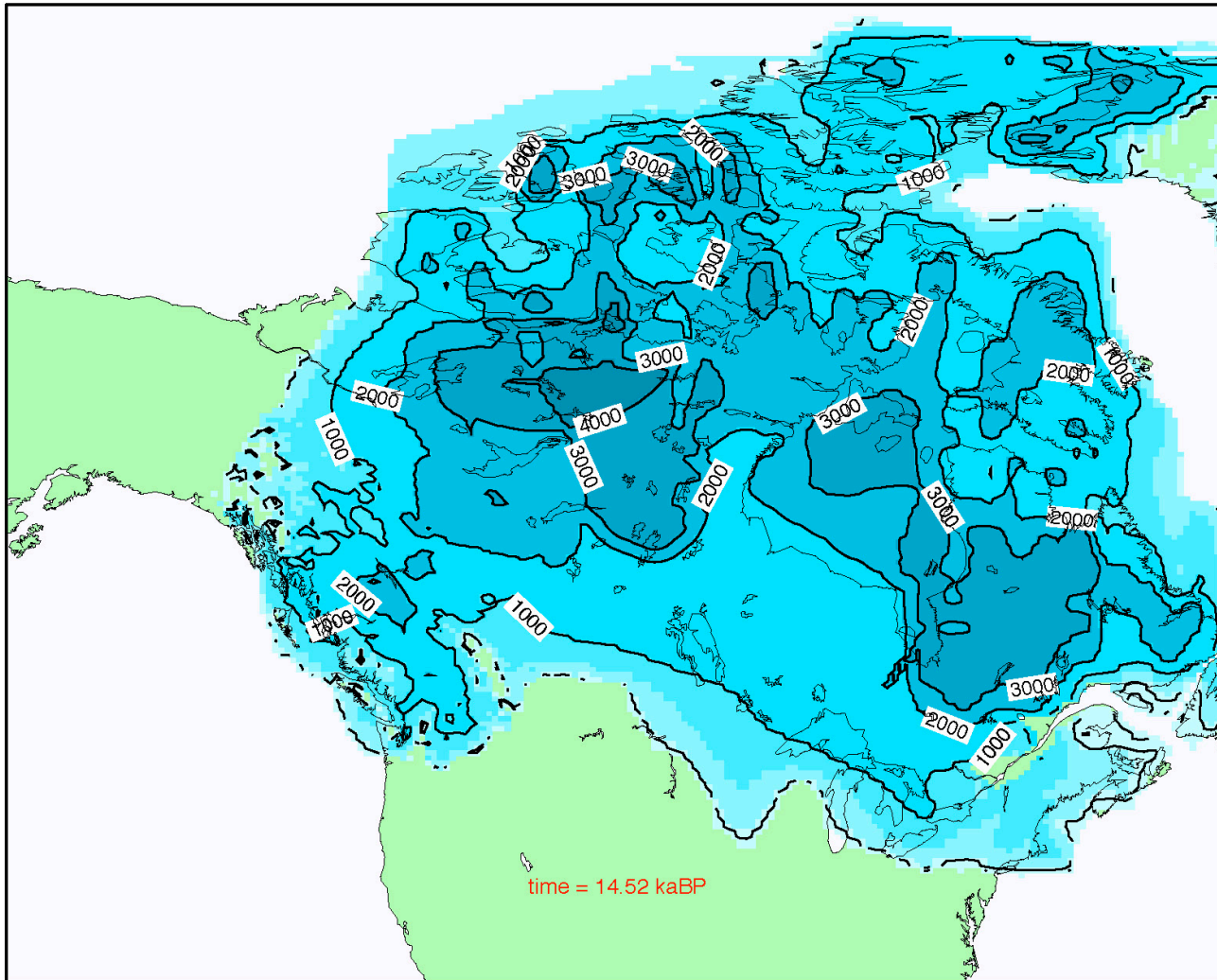
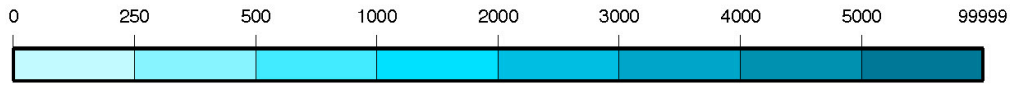
38_6_4_20_16_12.d5K#14489F



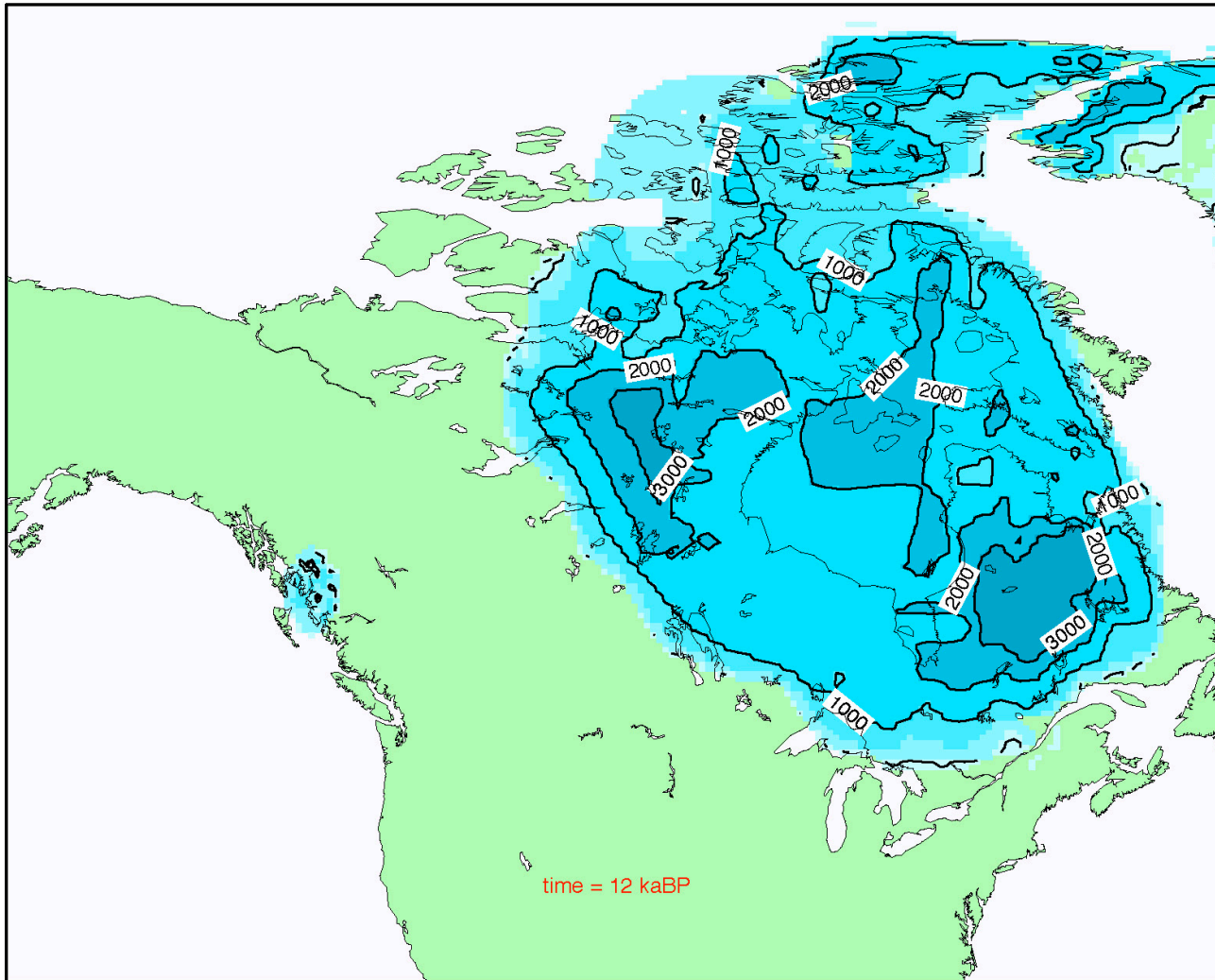
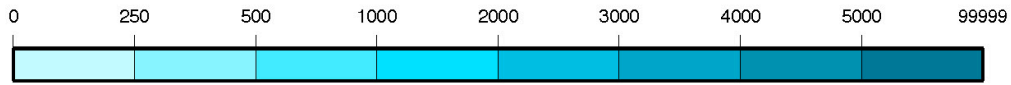
Litho- sphere 80 km	Upper mantle (4-5) $\times 10^{20}$ Pa s	Lower mantle (5-20) $\times 10^{22}$ Pa s
---------------------------	---	--



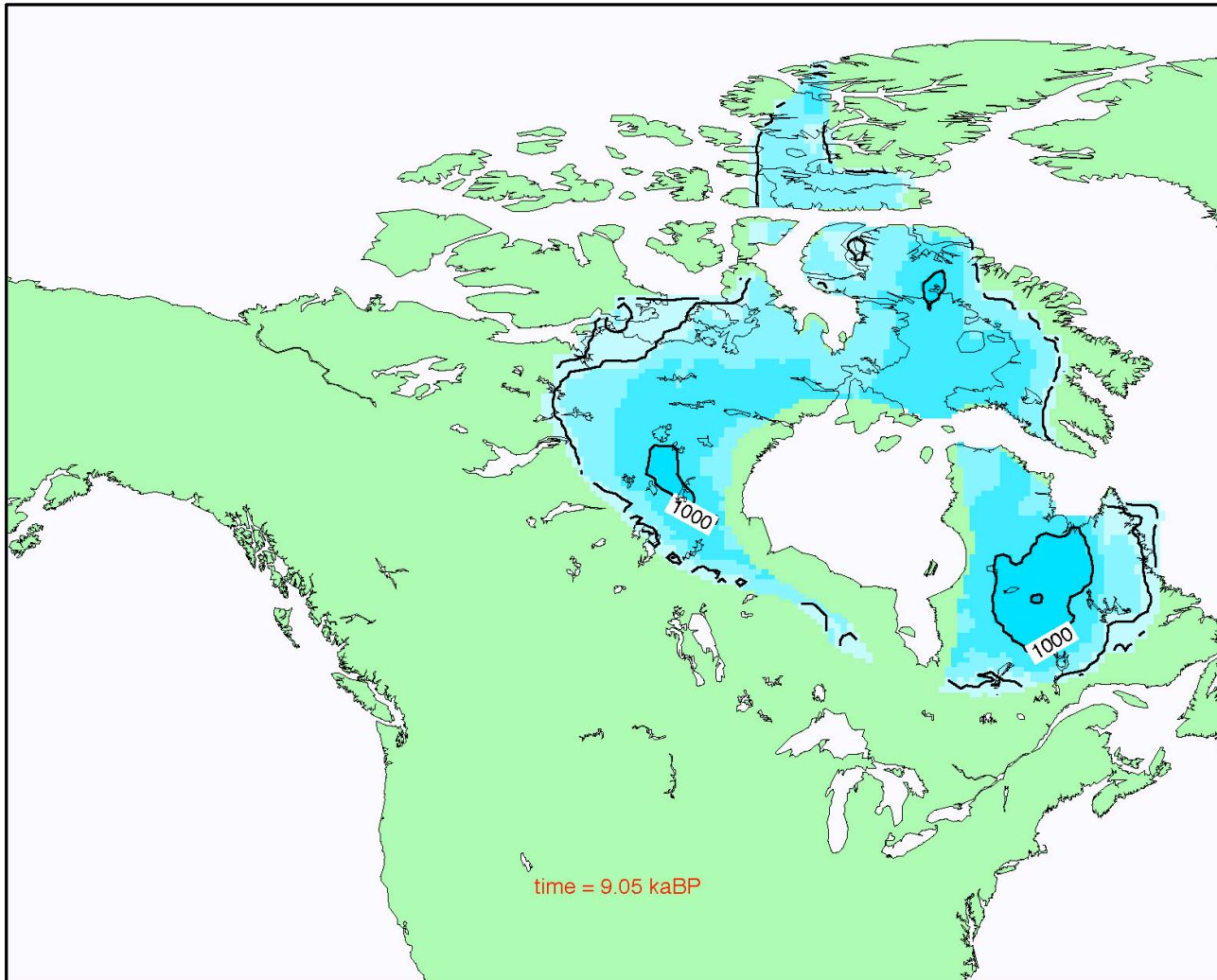
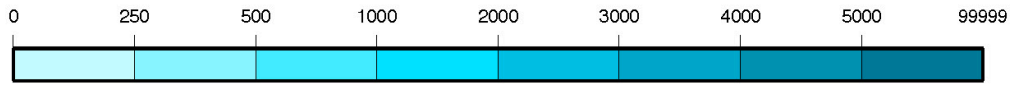
LGM 20 ka



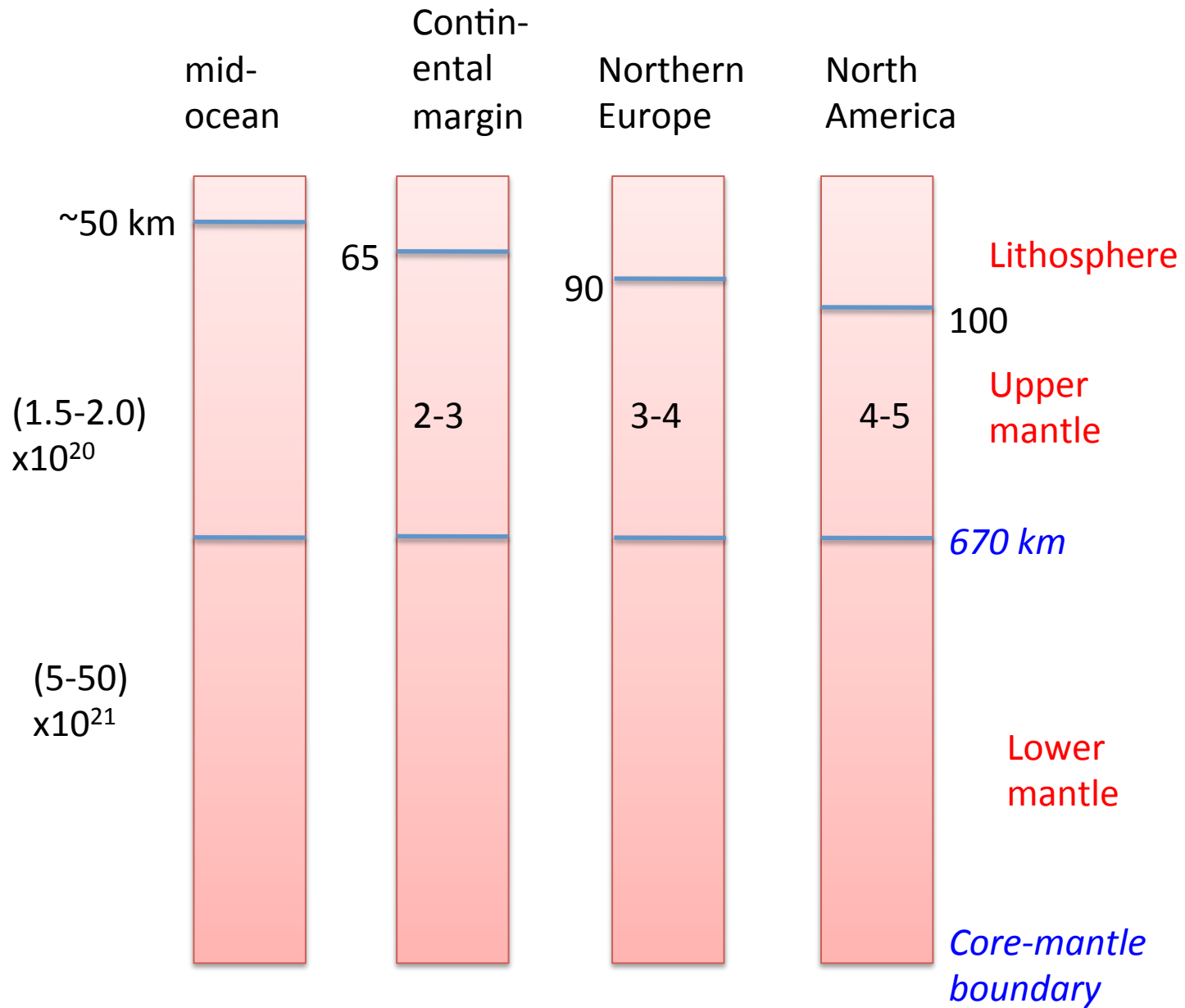
Onset of BA
14.5 ka

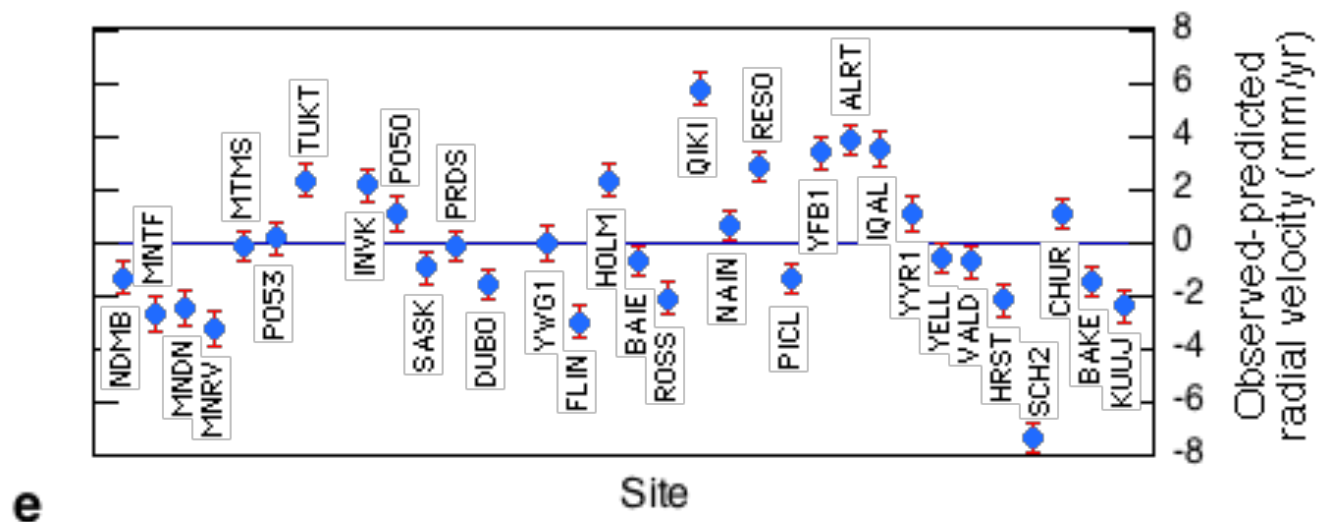
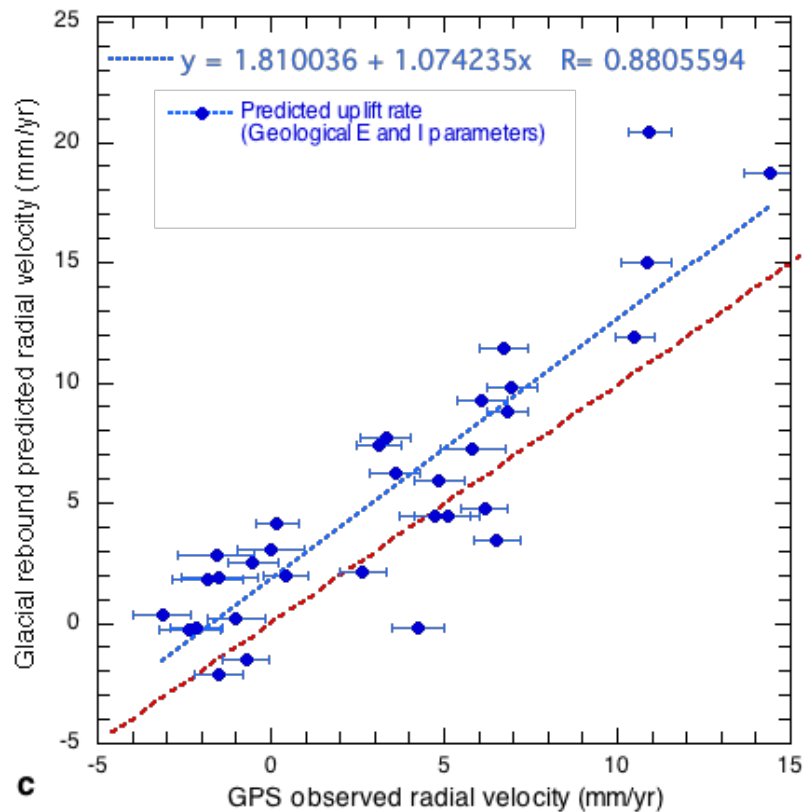
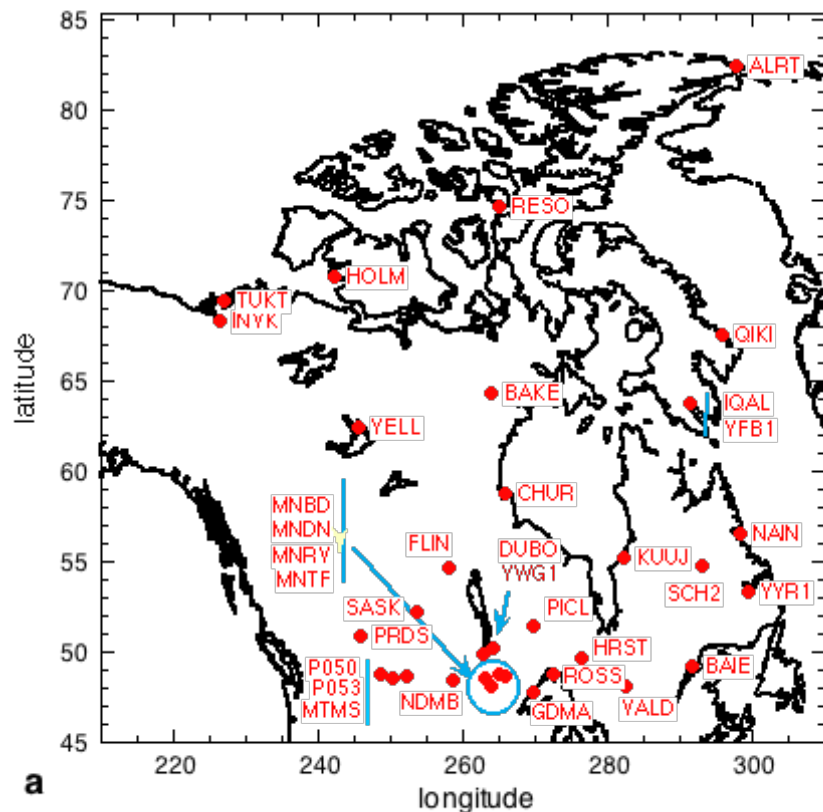


YD 12ka



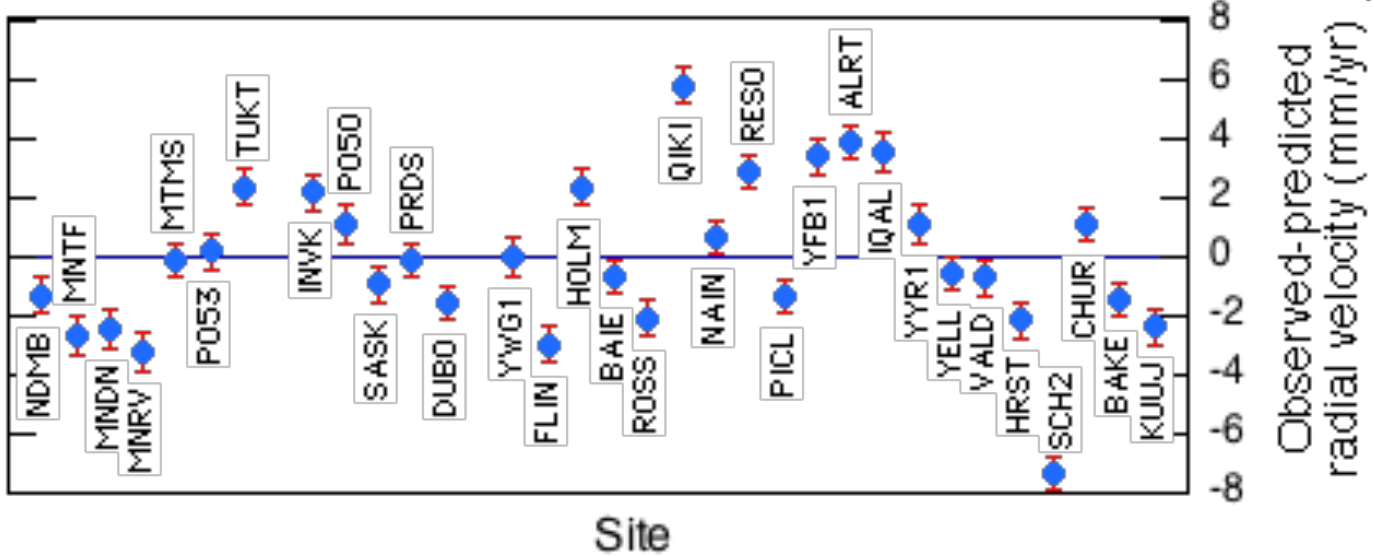
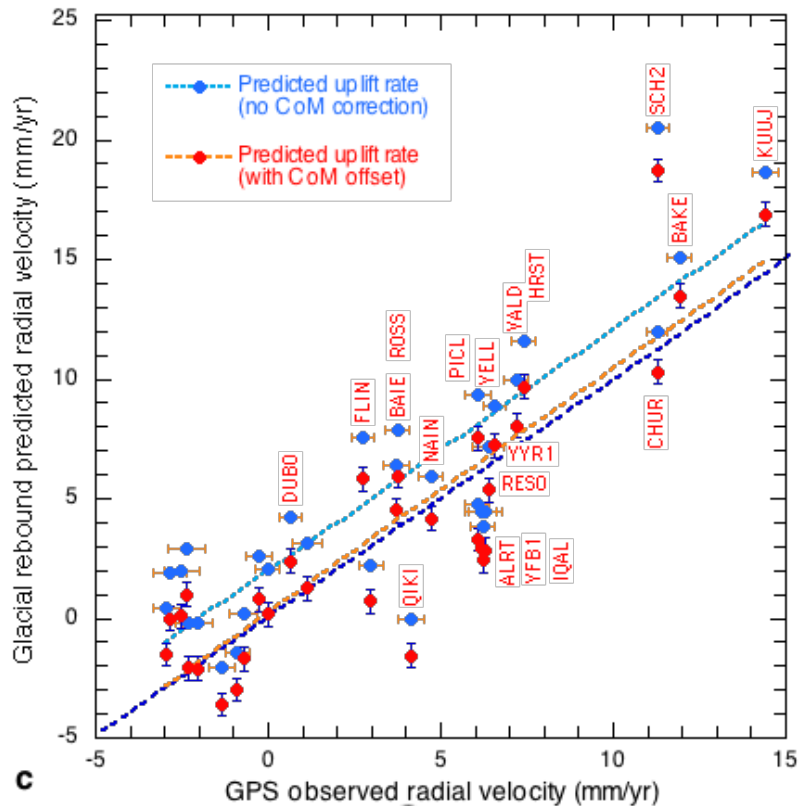
9.0 ka





$$U_z = \langle (u_{\text{obs}}^i - u_{\text{iso-pred}}^i) \sin \phi^i \rangle_W$$

$$= -1.62 \pm 0.51 \text{ mm/year}$$



	H (km)	η_{um} ($\times 10^{20}$ Pa s)	η_{lm} ($\times 10^{21}$ Pa s)
Geological data ^{1,2}	80-90	3.0-4.0	7-30
Tide gauge data ^{1,3}	80-100	4.0-4.5	≥ 10
GPS data ⁴	93-110	3.4-5.0	7-13

¹Lambeck and Purcell (2003); ²Lambeck et al., (2008); ³Lambeck and Ekman (1998); ⁴Zhao et al. (2012).

• North America	88 m
• Scandinavia, including Barents Sea	19 m
• Greenland	3 m
• British Isles	1 m
• Alpine glaciers	4 m
– Alaska	0.8 m
– South America	0.9 m
– Sub-Antarctic Islands	0.15 m
– Alps	0.07m
– Northern Asia (Taymyr, Putorama, Severnaya Z)	1.9 m
– Tibet and other mountain glaciers	1.0 m
<hr/>	
• Total	115 m
• Observed	134 m
	<hr/>
	'Missing Ice'
	19 m